

# Asymptotic limits of the Shallow Water equations

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Work in collaboration with:

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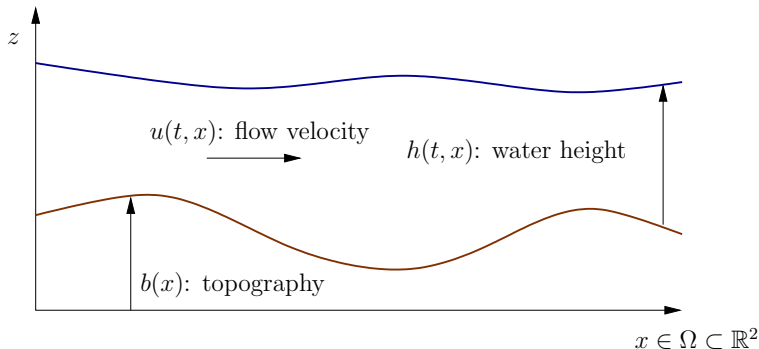


# Shallow Water equations

Shallow Water equations:

$$\partial_t h + \operatorname{div}(hu) = 0,$$

$$\partial_t(hu) + \operatorname{div}(hu \otimes u) + gh\nabla h = -gh\nabla b.$$

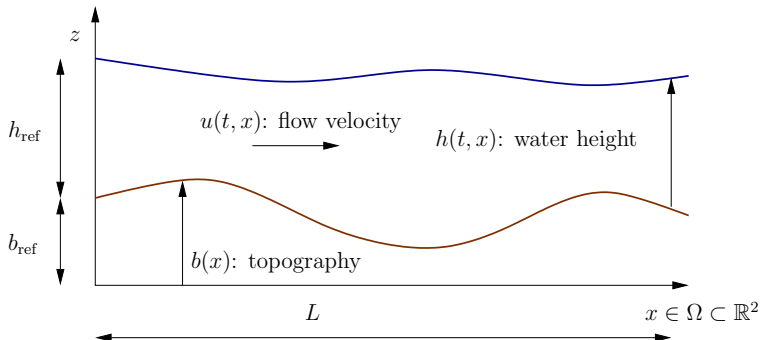


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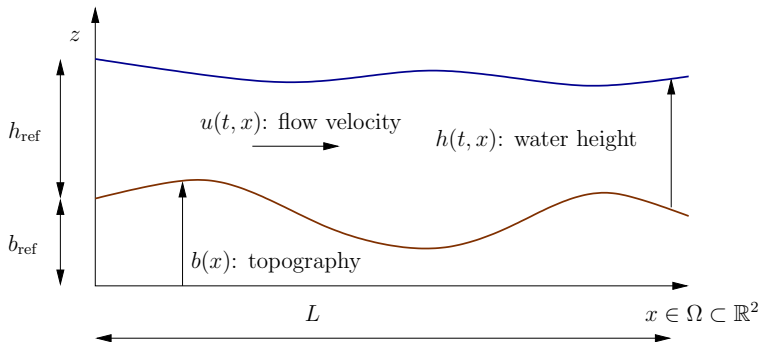


# Shallow Water equations

Dimensionless Shallow Water equations:

$$\frac{L}{t_{\text{ref}} u_{\text{ref}}} \partial_t h + \text{div}(hu) = 0,$$

$$\frac{L}{t_{\text{ref}} u_{\text{ref}}} \partial_t(hu) + \text{div}(hu \otimes u) + \frac{gh_{\text{ref}}}{u_{\text{ref}}^2} h \nabla h = -\frac{gh_{\text{ref}}}{u_{\text{ref}}^2} \frac{b_{\text{ref}}}{h_{\text{ref}}} h \nabla b.$$



# Shallow Water equations

Dimensionless Shallow Water equations:

$$Sr \partial_t h + \operatorname{div}(hu) = 0,$$

$$Sr \partial_t(hu) + \operatorname{div}(hu \otimes u) + \frac{1}{Fr^2} h \nabla h = -\frac{1}{Fr^2} \beta h \nabla b.$$

with

$$Sr(= St) = \frac{L}{t_{\text{ref}} u_{\text{ref}}} \text{ the Strouhal number (vortex),}$$

$$Fr = \frac{u_{\text{ref}}}{\sqrt{g h_{\text{ref}}}} \text{ the Froude number (flow vs gravity waves velocities)}$$

$$\text{and } \beta = \frac{b_{\text{ref}}}{h_{\text{ref}}}.$$

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with

$$Sr(= St) = \frac{L}{t_{\text{ref}} u_{\text{ref}}} \text{ the Strouhal number,}$$

$$Fr = \frac{u_{\text{ref}}}{\sqrt{g h_{\text{ref}}}} = \varepsilon^\alpha \quad (\varepsilon \ll 1) \text{ the Froude number}$$

$$\text{and } \beta = \frac{b_{\text{ref}}}{h_{\text{ref}}} = 1.$$

# Multiple scales in Shallow Water equations

Low Froude number flows:

velocities of the flow  $<$  speed of the gravity waves

$\implies$  multiple length / time scales  
(depending on initial and boundary conditions).

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During  $t_{\text{ref}}$ :

$$t_{\text{ref}} u_{\text{ref}} = L / \text{Sr}$$

distance of an advected particle

$<$

$$t_{\text{ref}} \sqrt{gh_{\text{ref}}} = (L / \text{Sr}) / \varepsilon^\alpha$$

distance of gravity waves.



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$$t_{\text{ref}} \sqrt{g h_{\text{ref}}} = (L / \text{Sr}) / \varepsilon^\alpha$$

distance of gravity waves.

In a  $O(L)$  domain:

$$L / u_{\text{ref}} = \text{Sr} t_{\text{ref}}$$

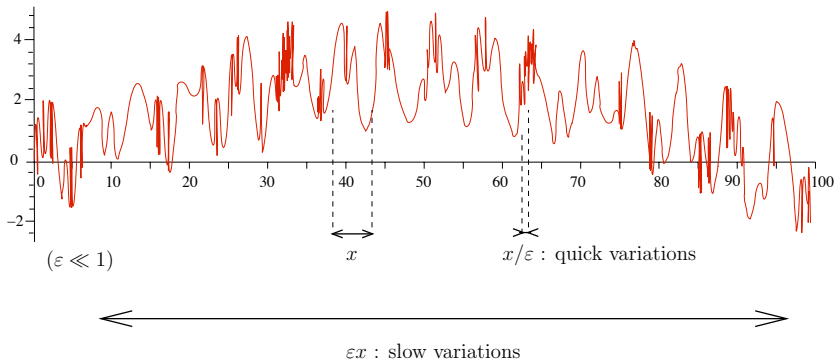
time scales for advected particle

$>$

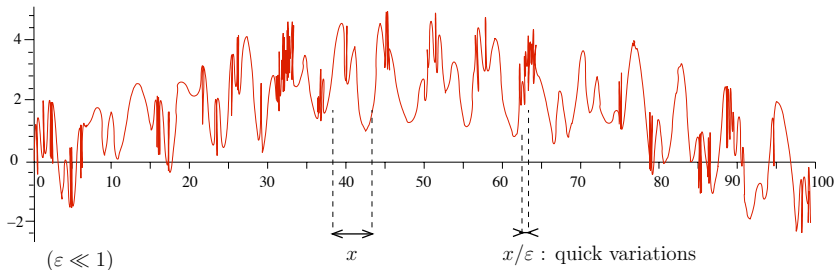
$$L / \sqrt{g h_{\text{ref}}} = \varepsilon^\alpha \text{Sr} t_{\text{ref}}$$

time scales for gravity waves.

# Multiscale topography



# Multiscale topography



$\varepsilon x$  : slow variations

$$X = \frac{x}{\varepsilon}$$

$$\chi = \varepsilon x.$$

# Outline

- 1 Balanced flow, topography at the 'normal' scale:  $b = b(x)$
- 2 Balanced flow, topography with quick variations:  $b = b(X, x)$ 
  - Weakly nonlinear regime
  - Fully nonlinear regime
- 3 Topography with long scale variations:  $b = b(x, \chi)$

## Formal derivations



D. Bresch, R. Klein, C. L., 2011

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$$b = b(x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

Flow on **advective time scales**.

Shallow Water equations:

$$\partial_t h + \operatorname{div}(hu) = 0,$$

$$\partial_t(hu) + \operatorname{div}(hu \otimes u) + \frac{1}{\varepsilon^2} h \nabla h = \frac{1}{\varepsilon^2} h \nabla b.$$

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Asymptotic development:

$$h(t, x, \varepsilon) = \sum_i \varepsilon^i h^i(t, x),$$

$$u(t, x, \varepsilon) = \sum_i \varepsilon^i u^i(t, x).$$

$$b = b(x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$O(\varepsilon^{-2})$$

$$h^0 \nabla (h^0 + b) = 0,$$

$$O(\varepsilon^{-1})$$

$$h^1 \nabla (h^0 + b) + h^0 \nabla h^1 = 0,$$

$$O(\varepsilon^0)$$

$$\partial_t h^0 + \text{div}(h^0 u^0) = 0,$$

$$\begin{aligned} \partial_t (h^0 u^0) + \text{div}(h^0 u^0 \otimes u^0) + \\ h^2 \nabla (h^0 + b) + h^1 \nabla h^1 + h^0 \nabla h^2 = 0. \end{aligned}$$



$$b = b(x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$O(\varepsilon^{-2})$$

$$h^0 \nabla (h^0 + b) = 0, \quad h^0 + b \equiv c^0(t)$$

$$O(\varepsilon^{-1})$$

$$h^1 \nabla (h^0 + b) + h^0 \nabla h^1 = 0, \quad h^1 \equiv c^1(t)$$

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$$\partial_t h^0 + \operatorname{div}(h^0 u^0) = 0,$$

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$$\frac{\partial b}{\partial t} = 0 : \quad \operatorname{div}(h^0 u^0) = -\frac{d}{dt} c^0(t), \quad \frac{dc^0}{dt} = -\frac{1}{|\Omega|} \int_{\Omega} h^0 u^0 \cdot n \, d\sigma$$

$$\partial_t (h^0 u^0) + \operatorname{div}(h^0 u^0 \otimes u^0) + h^0 \nabla h^2 = 0.$$

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Shallow Water limit when  $b = b(x)$ ,  $\text{Fr} = \varepsilon$ ,  $\text{Sr} = 1$ :

### Lake equations

$$\begin{aligned} \partial_t(h^0 u^0) + \text{div}(h^0 u^0 \otimes u^0) + h^0 \nabla h^2 &= 0, \\ \frac{dc^0}{dt} = \frac{dh^0}{dt} &= -\frac{1}{|\Omega|} \oint_{\Omega} h^0 u^0 \cdot n \, d\sigma \end{aligned}$$

+ initial / boundary conditions on  $h^0$ ,  $c^0$ .

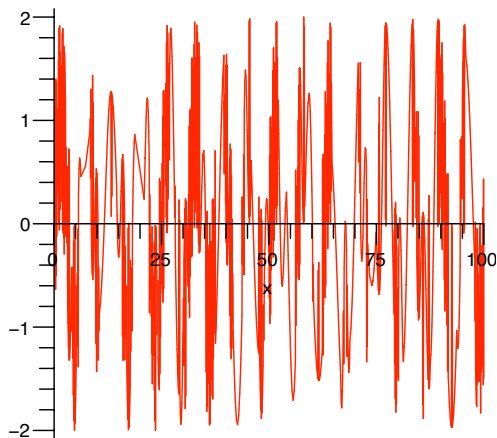


see D. Bresch, G. Métivier, AMRX, 2010  
for a rigorous justification of the limit.

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$$b = b(X, x)$$



$$X \in \mathbb{T}^2$$

$$b = b(X, x), \text{Fr} = \varepsilon^{3/2}, \text{Sr} = \varepsilon^{-1}$$

Characteristic lengths too short to support gravity waves.  
Weakly nonlinear regime.

Shallow Water equations:

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$$\partial_t(hu) + \varepsilon \operatorname{div}(hu \otimes u) + \frac{1}{\varepsilon^2} h \nabla h = \frac{1}{\varepsilon^2} h \nabla b.$$

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$$h^0 \nabla_X (h^0 + b) = 0,$$

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$$h^0 \nabla_x (h^0 + b) + h^1 \nabla_X (h^0 + b) + h^0 \nabla_X h^1 = 0,$$

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$$\overline{h^0 \nabla_x (h^0 + b)}^X + \overline{h^0 \nabla_X h^1}^X = \overline{h^0 \nabla_x (h^0 + b)}^X = 0 : c(t, x) = c(t).$$



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...

$$O(\varepsilon^0)$$

$$\partial_t h^0 + \operatorname{div}_X (h^0 u^0) = 0,$$

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...

$$O(\varepsilon^0)$$

$$\partial_t h^0 + \text{div}_X (h^0 u^0) = 0, \quad \overline{\partial_t (h^0 + b)}^X = 0 : c(t) = c,$$

$$h^0(X, x) = -b(X, x) + c.$$

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$$h^0 \nabla_X h^1 = 0 : \quad h^1(t, X, x) = h^1(t, x) \xrightarrow{O(\varepsilon^{-1})} h^1(t, x) = h^1(t).$$

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$$\int_{\Omega} \overline{O(\varepsilon^1)}^X dx : \quad \frac{dh^1}{dt} = -\frac{1}{|\Omega|} \oint_{\Omega} \overline{h^0 u^0}^X \cdot n \, d\sigma$$

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We assume rigid vertical walls on  $\partial\Omega$ :  $h^1 = cst = 0$ .

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$$\operatorname{div}_x \overline{(h^0 u^0)}^X = 0.$$

$$b = b(X, x), \text{Fr} = \varepsilon^{3/2}, \text{Sr} = \varepsilon^{-1}$$

Using each equation we obtain:

$$\begin{cases} \partial_t(h^0 u^0) + \operatorname{div}_X(h^0 u^0 \otimes u^0) + h^0 \nabla_x h^2 + h^0 \nabla_X h^3 = 0 \\ \operatorname{div}_X(h^0 u^0) = 0 \\ \operatorname{div}_x \overline{h^0 u^0}^X = 0 \\ \nabla_X h^2 = 0 \end{cases}$$

with  $h^0(X, x) = c - b(X, x)$ .



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☑ energy principle

with  $h^0(X, x) = c - b(X, x)$ .

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☑ energy principle

with  $h^0(X, x) = c - b(X, x)$ .

→ large scale ? Average in  $X$ .

→ small scale ?  $\tilde{h} = h - \bar{h}^X$ .

$$b = b(X, x), \text{Fr} = \varepsilon^{3/2}, \text{Sr} = \varepsilon^{-1}$$

At large scale:

$$\begin{cases} \partial_t (\overline{h^0 u^0})^X + \overline{h^0}^X \nabla_x h^2 = -\overline{h^3 \nabla_X b}^X \\ \text{div}_x \overline{h^0 u^0}^X = 0 \end{cases}$$

- response of the leading-order large-scale flow to accumulated small-scale pressure forces on the topography,
- the second order  $h^2$  acts like a Lagrangian multiplier.

$$b = b(X, x), \text{Fr} = \varepsilon^{3/2}, \text{Sr} = \varepsilon^{-1}$$

At small scale:

$$\begin{cases} \partial_t(\widetilde{h^0 u^0}) + \text{div}_X(h^0 u^0 \otimes u^0) + \widetilde{h^0 \nabla_X h^3} = \widetilde{b} \nabla_x h^2 \\ \text{div}_X(\widetilde{h^0 u^0}) = 0 \end{cases}$$

- interactions between small and large scales,
- $\nabla_x h^2$  acts on the fluctuations of the topography to drive the small scale flow.

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Weakly nonlinear limit version of the lake equations with oscillatory topography.



D. Bresch, D. Gérard-Varet, AML, 2007

$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

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Asymptotic development:

$$h(t, x, \varepsilon) = \sum_i \varepsilon^i h^i(t, X, x),$$

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$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$O(\varepsilon^{-3}) \\ h^0 \nabla_X (h^0 + b) = 0,$$



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$$O(\varepsilon^{-3}) \quad h^0 \nabla_X (h^0 + b) = 0, \quad h^0(t, X, x) + b(X, x) \equiv c(t, x)$$

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$$O(\varepsilon^{-2}) \quad h^0 \nabla_x (h^0 + b) + h^1 \nabla_X (h^0 + b) + h^0 \nabla_X h^1 = 0,$$

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$$O(\varepsilon^{-1}) \quad \text{div}_X (h^0 u^0) = 0, \\ \text{div}_X (h^0 u^0 \otimes u^0) + h^0 \nabla_x h^1 + h^0 \nabla_X h^2 = 0.$$

→ small scale ?

→ large scale ?

$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$u^0 \cdot \nabla_X u^0 + \nabla_X h^2 = -\nabla_x h^1.$$

Small scale:

$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$u^0 \cdot \nabla_X u^0 + \nabla_X h^2 = -\nabla_x h^1.$$

Small scale:

Taking the curl ( $\zeta = \text{curl} u = -\partial_{X_2} u_1 + \partial_{X_1} u_2$ ):

$$u^0 \cdot \nabla_X \zeta^0 + \zeta^0 \text{div}_X u^0 = \text{div}_X (\zeta^0 u^0) = 0,$$

as  $\text{div}_X (h^0 u^0) = 0$ , it reads  $h^0 u^0 \cdot \nabla_X (\zeta^0 / h^0) = 0$ .

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as  $\text{div}_X (h^0 u^0) = 0$ , it reads  $h^0 u^0 \cdot \nabla_X (\zeta^0 / h^0) = 0$ .

$$\zeta^0 = H^0 Q(\psi^{*,0}, x, t),$$

if  $Q$  is a **potential vorticity distribution function**,

if  $\psi^{*,0}$  is a **stream function** for  $h^0 u^0$ ,

$$\psi^{*,0} = \psi^0 + X^\perp \cdot \overline{h^0 u^0}^X \text{ with } h^0 u^0 = \nabla_X^\perp \psi^{*,0},$$

$$h^0 \nabla_X^2 \psi^0 - \nabla_X h^0 \cdot \nabla_X \psi^0 = (h^0)^3 Q(\psi^{*,0}, x, t) - \nabla_X h^0 \cdot \overline{h^0 u^0}^{X^\perp}.$$

Cell problem for a stationary vortical flow over variable topography,

$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$u^0 \cdot \nabla_X u^0 + \nabla_X h^2 = -\nabla_x h^1.$$

Small scale:

Taking the curl ( $\zeta = \text{curl} u = -\partial_{X_2} u_1 + \partial_{X_1} u_2$ ):

$$u^0 \cdot \nabla_X \zeta^0 + \zeta^0 \text{div}_X u^0 = \text{div}_X (\zeta^0 u^0) = 0,$$

as  $\text{div}_X (h^0 u^0) = 0$ , it reads  $h^0 u^0 \cdot \nabla_X (\zeta^0 / h^0) = 0$ .

$$\zeta^0 = H^0 Q(\psi^{*,0}, x, t),$$

if  $Q$  is a **potential vorticity distribution function**,

if  $\psi^{*,0}$  is a **stream function** for  $h^0 u^0$ ,

$$\psi^{*,0} = \psi^0 + X^\perp \cdot \overline{h^0 u^0}^X \text{ with } h^0 u^0 = \nabla_X^\perp \psi^{*,0},$$

$$h^0 \nabla_X^2 \psi^0 - \nabla_X h^0 \cdot \nabla_X \psi^0 = (h^0)^3 Q(\psi^{*,0}, x, t) - \nabla_X h^0 \cdot \boxed{\overline{h^0 u^0}^{X^\perp}}.$$

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Large scale:

$$b = b(X, x), \text{Fr} = \varepsilon, \text{Sr} = 1$$

$$u^0 \cdot \nabla_X u^0 + \nabla_X h^2 = -\nabla_x h^1.$$

Large scale (average in  $X$ ):

$$U \cdot T + \nabla_x h^1 = -q,$$

with

$$U = \overline{h^0 u^0}^X, \quad \tilde{u} = u^0 - \frac{1}{h^0} \overline{h^0 u^0}^X = \frac{1}{h^0} \nabla_X^\perp \psi^0,$$

$$T = \overline{\frac{1}{h^0} \nabla_X \tilde{u}}^X \quad \text{and} \quad q = \overline{\tilde{u} \cdot \nabla_X \tilde{u}}^X.$$

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Large scale (average in  $X$ ):

$$U \cdot T + \nabla_x h^1 = -q, \quad \text{Darcy type problem}$$

with

$$U = \overline{h^0 u^0}^X, \quad \tilde{u} = u^0 - \frac{1}{h^0} \overline{h^0 u^0}^X = \frac{1}{h^0} \nabla_X^\perp \psi^0,$$

$$T = \overline{\frac{1}{h^0} \nabla_X \tilde{u}}^X \quad \text{and} \quad q = \overline{\tilde{u} \cdot \nabla_X \tilde{u}}^X \quad (\text{small scale viscous forces}).$$

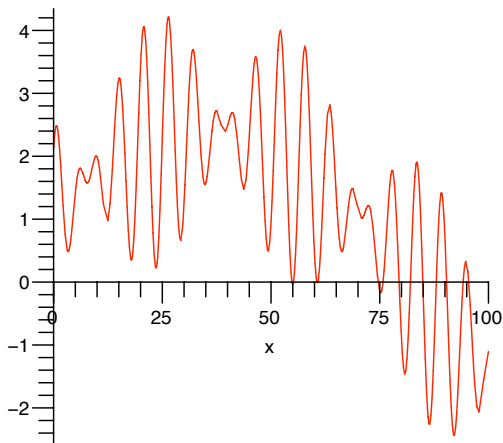
$\overline{O(\varepsilon^0)}^X$ :

$$\text{div}_x U = -\text{div}_x ((\nabla_x h^1 + q) \cdot T^{-1}) = -\frac{d\overline{h^0}^X}{dt}.$$

# Outline

- 1 Balanced flow, topography at the 'normal' scale:  $b = b(x)$
- 2 Balanced flow, topography with quick variations:  $b = b(X, x)$ 
  - Weakly nonlinear regime
  - Fully nonlinear regime
- 3 Topography with long scale variations:  $b = b(x, \chi)$

$$b = b(x, \chi), \text{Fr} = \varepsilon, \text{Sr} = 1$$



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Advective times for the normal scale  $L$ ,  
with gravity wave dynamics on a large scale  $L/\varepsilon$ .

Shallow Water equations:

$$\partial_t h + \operatorname{div}(hu) = 0,$$

$$\partial_t(hu) + \operatorname{div}(hu \otimes u) + \frac{1}{\varepsilon^2} h \nabla h = \frac{1}{\varepsilon^2} h \nabla b.$$

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Asymptotic development:

$$h(t, x, \varepsilon) = \sum_i \varepsilon^i h^i(t, x, \chi),$$

$$u(t, x, \varepsilon) = \sum_i \varepsilon^i u^i(t, x, \chi).$$

$$b = b(x, \chi), \text{Fr} = \varepsilon, \text{Sr} = 1$$

We get:

- $h^0 + b = c = c(t, x, \chi)$
- $\text{div}_x(h^0 u^0) = 0$
- $h^1 = h^1(t, \chi)$
- $$\begin{cases} \partial_t(h^0 u^0) + \text{div}_x(h^0 u^0 \otimes u^0) + h^0 \nabla_x h^2 + h^0 \nabla_\chi h^1 = 0, \\ \partial_t h^1 + \text{div}_x(h^0 u^1) + \text{div}_x(h^1 u^0) + \text{div}_\chi(h^0 u^0) = 0. \end{cases}$$



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—→ average in  $x$ : long-wave equations

—→ study of the small scale flow

$$b = b(x, \chi), \text{Fr} = \varepsilon, \text{Sr} = 1$$

Long wave:

$$\begin{cases} \partial_t (\overline{h^0 u^{0^x}}) + \overline{h^{0^x}} \nabla_\chi h^1 = \overline{h^2 \nabla_x h^{0^x}} \\ \partial_t h^1 + \text{div}_\chi (\overline{h^0 u^{0^x}}) = 0. \end{cases}$$

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$\approx$  standard linearized shallow water equations

$\overline{h^2 \nabla_x h^{0^x}}$  (from  $h^0 \nabla_x h^2$ ): net resistance  
(small-scale flow through the rough topography).

$$b = b(x, \chi), \text{Fr} = \varepsilon, \text{Sr} = 1$$

Small scale:

$$\begin{cases} \partial_t \widetilde{\widetilde{h^0 u^0}} + \text{div}_x(h^0 u^0 \otimes u^0) + h^0 \nabla_x h^2 = -\widetilde{\widetilde{h^0}} \nabla_\chi h^1, \\ \text{div}_x \widetilde{\widetilde{h^0 u^0}} = 0, \end{cases}$$

where  $\varphi = \overline{\varphi}^x + \widetilde{\varphi}$ .

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- divergence free,
- $h^2$ : Lagrangian multiplier,
- small-scale flow driven by the long-wave unbalanced part of the large-scale height gradient.



R. Klein, JCP, 1995 (low Mach number)

$$b = b(x, \chi), \text{Fr} = \varepsilon, \text{Sr} = 1$$

If  $b(x, \chi) = b(\chi)$ :

$\implies$  wave equation with spatially varying signal speed for  $h^1$ :

$$\partial_t^2 h^1 - \operatorname{div}_\chi ((c - b(\chi)) \nabla_\chi h^1) = 0, \quad (1)$$

where  $c = b + h^0 \equiv \text{const.}$

## Concluding remarks

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 $\text{Fr} = \varepsilon$ ,  $\text{Sr} = 1$ : [Lake equations](#).
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 $\text{Fr} = \varepsilon^{3/2}$ ,  $\text{Sr} = \varepsilon^{-1}$ : The large-scale accumulation of net pressure forces drives the large-scale balanced flow; the large-scale height gradients produce small-scale forces driving the small-scale flow.
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- ③ Topography with long scale variations:  $b = b(x, \chi)$   
 $\text{Fr} = \varepsilon$ ,  $\text{Sr} = 1$ : as for the weakly nonlinear case, but the large-scale flow involves non-balanced terms.